



# **INVERSE MODELING ON SOIL SOLUTION WATER TO DETERMINE SOIL CHEMICAL COMPOSITION**

BRANDON MONTGOMERY

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# OVERVIEW

Introduction

Experiment/discussion

Conclusion

Questions



# RESEARCH QUESTION

Can PHREEQC be used to determine the chemical composition of soil based on soil solution water?



# Controls on Soil Solution Chemistry in a Subalpine Forest in North-Central Colorado

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## ABSTRACT

High-elevation ecosystems in the western USA are potentially susceptible to increased inputs of strong acids. A long-term research project was established to identify the processes controlling surface water chemistry and to evaluate the sensitivity of Loch Vale Watershed in Rocky Mountain National Park, Colorado, to acid precipitation. Using lysimeters, we estimated the concentration and flux of major solutes in the Oie and B horizons in an old-grown Engelmann spruce (*Picea engelmannii* Parry ex Engelm.) and subalpine fir [*Abies lasiocarpa* (Hook.) Nutt.] forest, and in an adjacent site disturbed by a snow avalanche. In the forested site, most solutes were highly concentrated in soil solutions during the initial stages of snowmelt, and concentrations declined rapidly to low levels in the first 4 to 6 wk of snowmelt. Surface water chemistry in Loch Vale Watershed is controlled principally by mineral weathering. During the early stages of snowmelt, however, the flushing of solutes, especially N, S, and C, from forest soils (which comprise only 6% of the study area) exerts an important influence. Rates of cationic denudation per unit area were 3.5 times higher in the forest ( $131 \mu\text{mol}_c \text{ m}^{-2} \text{ yr}^{-1}$ ) than in the whole watershed ( $38 \mu\text{mol}_c \text{ m}^{-2} \text{ yr}^{-1}$ ) probably because of  $\text{H}^+$  exudation from roots during nutrient uptake and the generation of organic acids in the forest soils. Rates of N mineralization and nitrification as well as concentrations of  $\text{NO}_3^-$  in the soil solution were higher in the second year after forest disturbance from a snow avalanche than immediately following the disturbance, indicating a delayed nitrification response. Significant amounts of  $\text{NH}_4^+$  and  $\text{NO}_3^-$  were temporarily stored in extractable forms in the soil during the first year after the disturbance. Although precipitation inputs and theoretical mineral weathering can explain the total annual solute flux from Loch Vale Watershed, the effects of forest soil solutes may be important during the initial stages of snowmelt and following large-scale disturbance.

term changes in soil solution chemistry during snowmelt may also result from the flushing of byproducts of decomposition and mineral weathering (Fahey and Yavitt, 1988) that could neutralize acidity. It is likely that a combination of natural biotic and abiotic factors and anthropogenic inputs are responsible for seasonal and long-term changes in surface water chemistry.

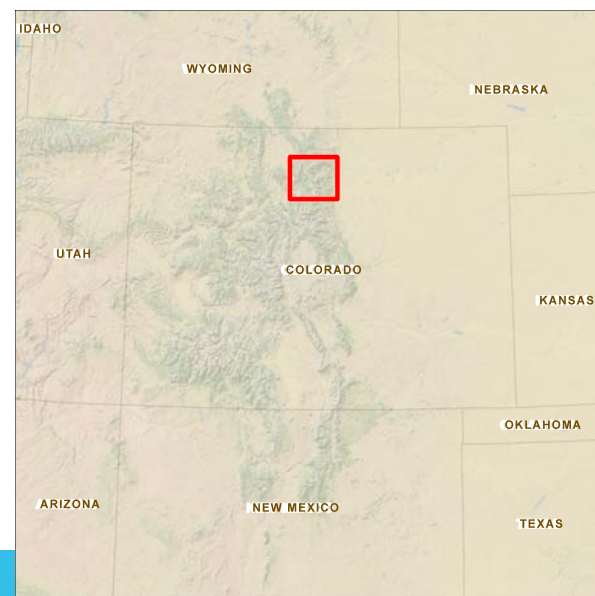
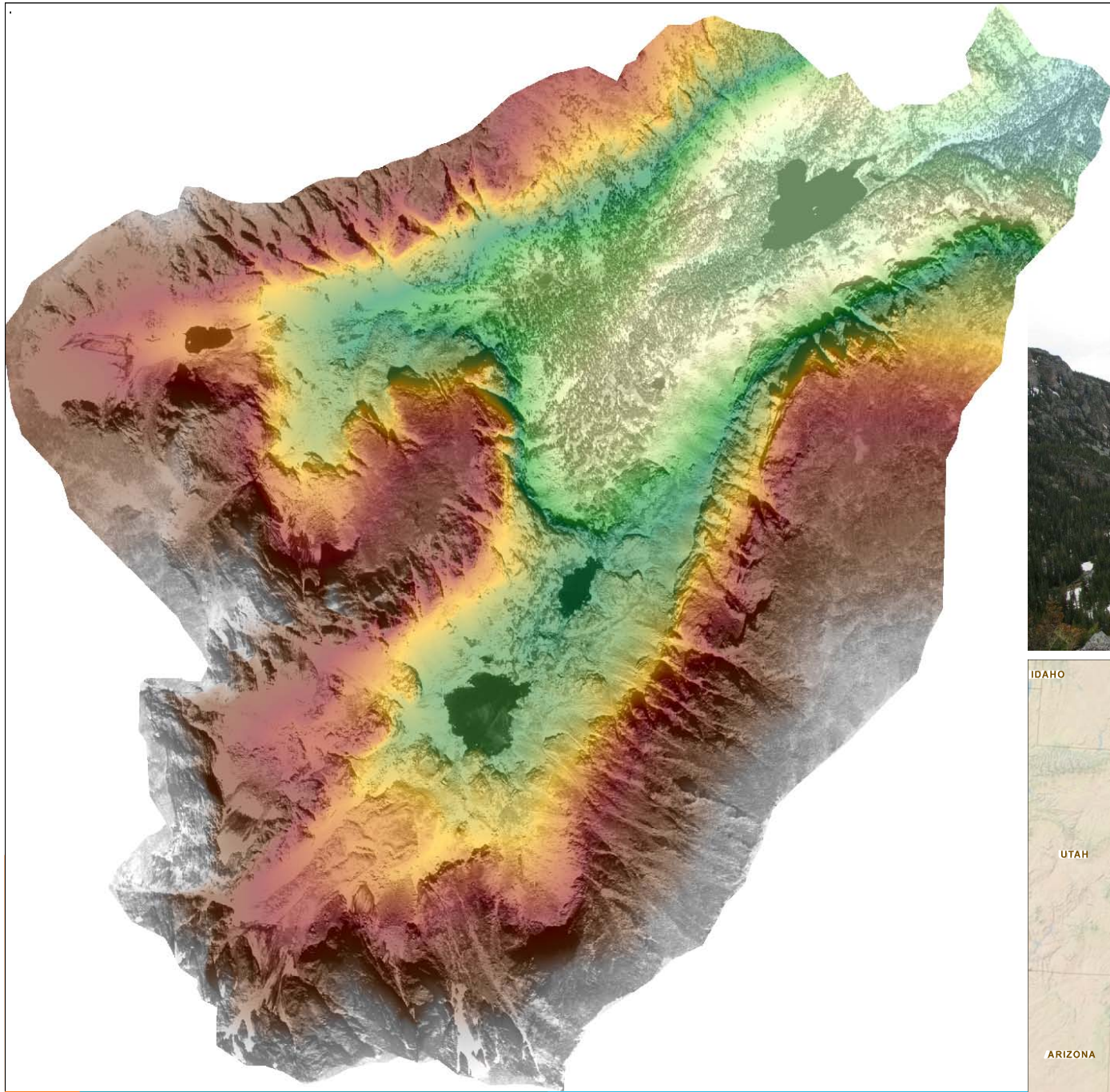
At Loch Vale Watershed, located in Rocky Mountain National Park, yearly average surface water pH is 6.3, but during spring snowmelt the pH of streamwater can be as low as 5.7 (Baron and Bricker, 1987). A long-term research project was established in 1983 to identify the processes controlling surface water chemistry and to evaluate the sensitivity of this watershed to acidic precipitation. Researchers have characterized the sensitivity of the soil environment to acid deposition (Walthall, 1985) and quantified rates of cationic denudation at the study area (Mast, 1989). This study of soil solution chemistry is part of a larger research project designed to quantify internal transfers of acidity and to evaluate the importance of the vegetation in influencing the geochemistry of the watershed (Arthur, 1990). The objectives of our study were to: (i) determine the importance of several mechanisms regulating soil solution fluxes, including adsorption by soil surfaces and decomposition and leaching in the forest floor, (ii) examine changes in soil solution chemistry with depth and time, and (iii) quantify element fluxes so that nutrient budgets could be calculated. Our hypotheses were that (i) concentrations of plant nutrients, especially  $\text{NH}_4^+$ ,  $\text{NO}_3^-$ , and K, would

# THE LOCH VALE WATERSHED



Photo Courtesy of Jake Tyson



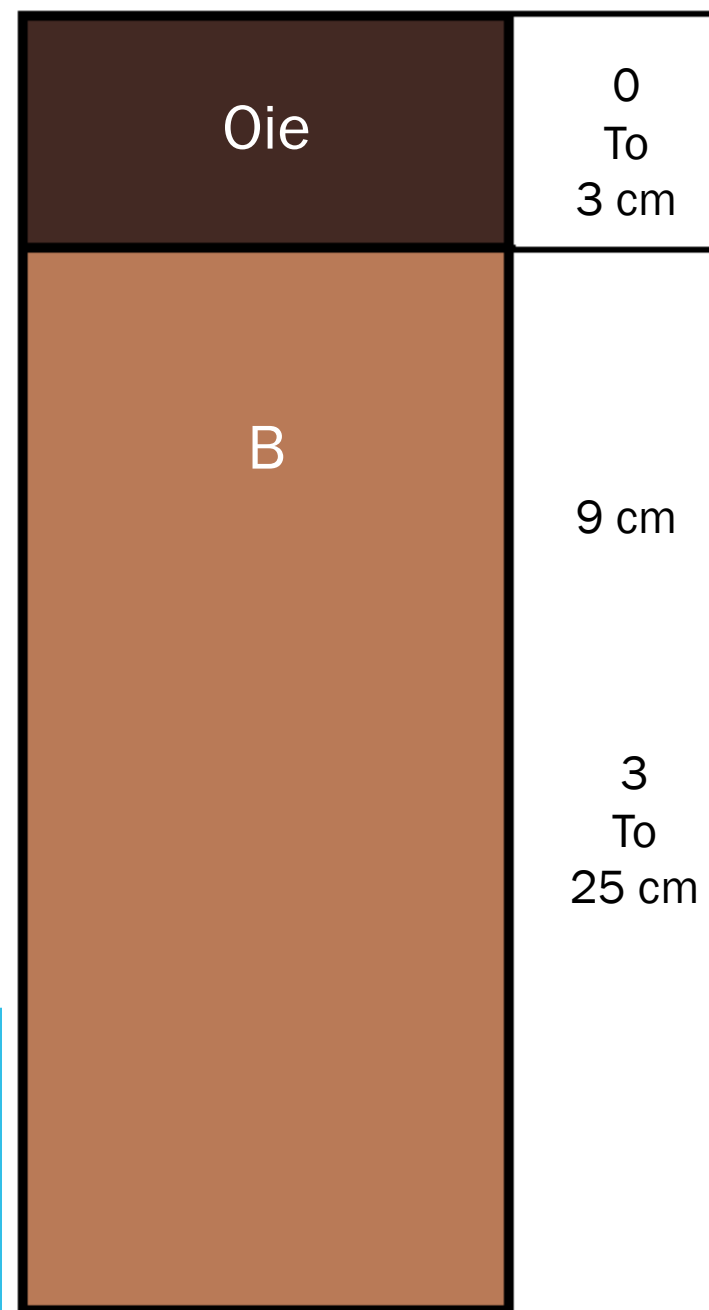


"Logistical support and/or data were provided by the USGS supported Loch Vale Watershed Long-term Ecological Research and Monitoring program and Colorado State University."

Solute	Zero-tension†		Tension‡
	O horizon	Rooting zone	Rooting zone
	$\mu\text{mol L}^{-1}$		
H	12 (131%)	8 (64%)	10 (44%)
Cl	55 (86%)	38 (71%)	33 (59%)
SO <sub>4</sub>	15 (62%)	16 (51%)	24 (34%)
Alkalinity	90 (83%)	72 (109%)	61 (69%)
Na	98 (102%)	81 (87%)	20 (50%)
K	70 (137%)	23 (87%)	24 (83%)
Mg	31 (88%)	30 (67%)	29 (39%)
Ca	73a (90%)	55a (54%)	138b (57%)
Si	36a (27%)	85b (75%)	85b (17%)
Al	13a (71%)	14a (51%)	31b (31%)
pH	4.9	5.0	5.0

† For zero-tension lysimeters, values are volume-weighted means for 3 yr (1986–1988). Rooted zone zero-tension lysimeters have mean depth of 9 cm.

‡ For tension lysimeters (9–28 cm), values are mean concentrations for 2 yr (1987–1988).



Lysimeter- used to collect leachate in soils, or provide measurements for evapotranspiration.



# Snowpack Melt and Precipitation Percolating through the Oie Horizon





## Input Waters

### SOLUTION 1

```
temp      1
pH        5.38
pe        4
redox     pe
units     mg/kgw
density   1
Ca        0.2842
Mg        0.04
Na        0.1828
K         0.08
S(6)      0.5157
Cl        0.1585
Si        0.60667
Al        10.2 ug/kgw
-water    1 # kg
```

### SOLUTION 3

```
temp      5
pH        5.48
pe        4
redox     pe
units     mg/kgw
density   1
Ca        0.29
Mg        0.031
K         0.02
Na        0.061
Cl        0.09
S(6)      0.81
-water    1 # kg
```

## Output Water

### SOLUTION 2

```
temp      5
pH        4.9
pe        4
redox     pe
units     umol/kgw
density   1
Cl        55
S(6)      15
Alkalinity 90
Na        98
K         70
Mg        31
Ca        73
Si        36
Al        13
-water    1 # kg
```

## Modeling Set-up

### INVERSE\_MODELING 1

```
-solutions      1      3      2
-uncertainty    0.75    0.75    0.75
-phases
  CO2(g)        force dis
  CaX2
  H2(g)
  H2O(g)        force dis
  MgX2
  NH4X
  NaX
  O2(g)         force dis
  AlX3
  KX
  N2(g)
-balances
  Al            0.75    0.75    0.75
  Alkalinity    0.75    0.75    0.75
  Ca            0.75    0.75    0.75
  Mg            0.75    0.75    0.75
  Na            0.75    0.75    0.75
  Si            0.75    0.75    0.75
  S(6)          0.75    0.75    0.75
  Cl            1      1      1
  H(0)          0.75    0.75    0.75
  K             0.75    0.75    0.75
-range          1000
-minimal
-tolerance      1e-010
-force_solutions false true false
-mineral_water  true
-multiple_precision true
-mp_tolerance   1e-012
-censor_mp      1e-020
```

## Solution 2:

	Input		Delta		Input+Delta
pH	4.900e+000	+	0.000e+000	=	4.900e+000
Al	1.300e-005	+	0.000e+000	=	1.300e-005
Alkalinity	9.000e-005	+	1.974e-005	=	1.097e-004
C(-4)	0.000e+000	+	0.000e+000	=	0.000e+000
C(4)	4.161e-003	+	8.152e-004	=	4.976e-003
Ca	7.300e-005	+	-5.475e-005	=	1.825e-005
Cl	5.500e-005	+	-5.053e-005	=	4.471e-006
H(0)	0.000e+000	+	0.000e+000	=	0.000e+000
K	7.000e-005	+	-5.250e-005	=	1.750e-005
Mg	3.100e-005	+	-2.325e-005	=	7.750e-006
N(-3)	0.000e+000	+	0.000e+000	=	0.000e+000
N(0)	0.000e+000	+	0.000e+000	=	0.000e+000
N(3)	0.000e+000	+	0.000e+000	=	0.000e+000
N(5)	0.000e+000	+	0.000e+000	=	0.000e+000
Na	9.800e-005	+	-7.350e-005	=	2.450e-005
O(0)	0.000e+000	+	0.000e+000	=	0.000e+000
S(-2)	0.000e+000	+	0.000e+000	=	0.000e+000
S(6)	1.500e-005	+	-5.605e-006	=	9.395e-006
Si	3.600e-005	+	-2.590e-005	=	1.010e-005
X	0.000e+000	+	0.000e+000	=	0.000e+000

Delta- Any adjustments to the original data (must fall within the uncertainty limits)

Solution fractions:			Minimum	Maximum
Solution	1	1.000e+000	5.093e-001	1.000e+000
Solution	3	0.000e+000	0.000e+000	4.908e-001
Solution	2	1.000e+000	1.000e+000	1.000e+000

Phase mole transfers:			Minimum	Maximum	
CO2(g)	4.976e-003	2.904e-003	7.281e-003	CO2	
CaX2	1.586e-005	6.797e-006	5.722e-005	CaX2	
H2(g)	1.707e-004	1.139e-004	2.724e+001	H2	
H2O(g)	0.000e+000	0.000e+000	2.724e+001	H2O	
MgX2	6.105e-006	4.870e-006	4.981e-005	MgX2	
NH4X	-1.138e-004	-1.621e-004	-7.591e-005	NH4X	
NaX	1.655e-005	1.058e-005	1.080e-004	NaX	
O2(g)	0.000e+000	0.000e+000	1.362e+001	O2	
AlX3	1.262e-005	2.588e-006	2.270e-005	AlX3	
KX	1.545e-005	1.392e-005	1.020e-004	KX	
N2(g)	5.690e-005	3.795e-005	8.103e-005	N2	

Redox mole transfers:	
H(0)	3.414e-004
N(-3)	-1.138e-004
N(0)	1.138e-004

Sum of residuals (epsilons in documentation):	8.814e+000
Sum of delta/uncertainty limit:	8.813e+000
Maximum fractional error in element concentration:	9.187e-001

- Precipitation of Ammonium
- Immobilization
- C:N Ratio

# Gravity Dependent Soil Solution Water Moving through the Mineral Horizon





□

Modeling Set-up

Input

Output

SOLUTION 2  
temp 5  
pH 4.9  
pe 4  
redox pe  
units umol/kgw  
density 1  
Cl 55  
S(6) 15  
Alkalinity 90  
Na 98  
K 70  
Mg 31  
Ca 73  
Si 36  
Al 13  
-water 1 # kg

SOLUTION 3  
temp 5  
pH 5  
pe 4  
redox pe  
units umol/kgw  
density 1  
S(6) 16  
Alkalinity 72  
Na 81  
K 23  
Mg 30  
Ca 55  
Si 85  
Al 14  
Cl 38  
-water 1 # kg

INVERSE\_MODELING 1  
-solutions 2 3  
-uncertainty 0.5 0.4  
-phases  
Al(OH)3(a)  
Albite dis  
Alunite dis  
Anhydrite  
Anorthite dis  
Aragonite dis  
Ca-Montmorillonite  
Calcite force  
Chalcedony  
Chlorite(14A)  
Chrysotile  
CO2(g) force  
Dolomite  
Gibbsite force dis  
Gypsum force  
H2(g)  
H2O(g)  
Halite  
Illite  
K-feldspar force dis  
K-mica force dis  
Kaolinite  
O2(g) dis  
Quartz force dis  
Sepiolite  
Sepiolite(d)  
SiO2(a)  
Talc dis  
-range 1000  
-minimal  
-tolerance 1e-010  
-mineral\_water true  
-multiple\_precision true  
-mp\_tolerance 1e-012  
-censor\_mp 1e-020

Phase mole transfers:		Minimum	Maximum	
Alunite	1.120e-005	0.000e+000	1.840e-005	
KA13(SO4)2(OH)6				
Ca-Montmorillon	-1.807e+001	-1.807e+001	-1.807e+001	
Ca0.165Al2.33Si3.67O10(OH)2				
Calcite	2.982e+000	2.981e+000	2.982e+000	CaCO3
Chrysotile	-9.685e+000	-9.685e+000	-9.685e+000	Mg3Si2O5(OH)4
CO2(g)	-2.978e+000	-2.978e+000	-2.978e+000	CO2
Gypsum	0.000e+000	-1.440e-005	2.240e-005	CaSO4:2H2O
Halite	4.860e-005	4.860e-005	5.320e-005	NaCl
K-mica	7.200e-006	0.000e+000	1.840e-005	
KA13Si3O10(OH)2				
Kaolinite	2.105e+001	2.105e+001	2.105e+001	Al2Si2O5(OH)4
Sepiolite	1.453e+001	1.453e+001	1.453e+001	
Mg2Si3O7.5OH:3H2O				

Redox mole transfers:

Sum of residuals (epsilons in documentation):	8.106e+000
Sum of delta/uncertainty limit:	8.106e+000
Maximum fractional error in element concentration:	4.000e-001

Phase mole transfers:		Minimum	Maximum	
Anhydrite	2.240e-005	-1.000e+003	1.000e+003	CaSO <sub>4</sub>
Anorthite	3.826e+000	0.000e+000	1.000e+003	CaAl <sub>2</sub> Si <sub>2</sub> O <sub>8</sub>
Ca-Montmorillon	-2.321e+001	-1.000e+003	1.000e+003	
Ca <sub>0.165</sub> Al <sub>2.33</sub> Si <sub>3.67</sub> O <sub>10</sub> (OH) <sub>2</sub>				
Calcite	3.576e-003	-1.000e+003	1.000e+003	CaCO <sub>3</sub>
Chlorite(14A)	-4.784e+000	-1.000e+003	1.000e+003	
Mg <sub>5</sub> Al <sub>2</sub> Si <sub>3</sub> O <sub>10</sub> (OH) <sub>8</sub>				
CO <sub>2</sub> (g)	0.000e+000	-1.000e+003	1.000e+003	CO <sub>2</sub>
Gibbsite	0.000e+000	0.000e+000	1.000e+003	Al(OH) <sub>3</sub>
Gypsum	0.000e+000	-1.000e+003	1.000e+003	CaSO <sub>4</sub> :2H <sub>2</sub> O
Halite	4.860e-005	4.860e-005	5.320e-005	NaCl
Illite	3.067e-005	-1.000e+003	3.067e-005	
K <sub>0.6</sub> Mg <sub>0.25</sub> Al <sub>2.3</sub> Si <sub>3.5</sub> O <sub>10</sub> (OH) <sub>2</sub>				
K-feldspar	0.000e+000	0.000e+000	1.000e+003	KAlSi <sub>3</sub> O <sub>8</sub>
K-mica	0.000e+000	0.000e+000	1.000e+003	
KAl <sub>3</sub> Si <sub>3</sub> O <sub>10</sub> (OH) <sub>2</sub>				
Kaolinite	2.800e+001	-1.000e+003	1.000e+003	Al <sub>2</sub> Si <sub>2</sub> O <sub>5</sub> (OH) <sub>4</sub>
Quartz	0.000e+000	0.000e+000	1.000e+003	SiO <sub>2</sub>
Sepiolite(d)	1.196e+001	-1.000e+003	1.000e+003	
Mg <sub>2</sub> Si <sub>3</sub> O <sub>7</sub> .5OH:3H <sub>2</sub> O				

## Redox mole transfers:

Sum of residuals (epsilons in documentation):	8.106e+000
Sum of delta/uncertainty limit:	8.106e+000
Maximum fractional error in element concentration:	4.000e-001

# CONCLUSIONS

Inverse modeling constraints

User constraints

Accuracy?





## SOURCES

Arthur, M.A., and T.J. Fahey. 1993. Controls on Soil Solution Chemistry in a Subalpine Forest in North-Central Colorado. Soil Sci. Soc. Am. J. 57: 1122-1130.

Rustad, Lindsey E., and C.S. Cronan. 1988. Element loss and retention during litter decay in a red spruce stand in Maine. Journal canadien de recherche forestiere. 18: 58-68.



**QUESTIONS?**

